

Sketch of an approach to replace the radiative transfer integrodifferential equations by a system of linear equations (see *Goody and Yung (GY)*, Chapters 2 and 8)

Expansion of azimuth dependence:

In general, scattering problems have azimuthal (ϕ) dependence, even though Φ may not be explicitly azimuthally-dependent, because of geometry:

$$\mu \frac{dI(\tau, \mu, \phi)}{d\tau} = I(\tau, \mu, \phi) - \frac{\omega(\tau)}{4\pi} \int d\Omega' \Phi(\tau, \mu, \phi, \mu', \phi') I(\tau, \mu', \phi') - \Sigma(\tau, \mu, \phi). \quad (\text{cf. } GY 8.1)$$

$\Sigma(\tau, \mu, \phi)$ is a “primary” source of radiation (*e.g.*, thermal; note that Goody and Yung treat the solar source separately).

- a. $\Phi(\tau, \mu, \phi, \mu', \phi')$ is expanded in *spherical harmonics*, $Y_{lm_l}(\theta, \phi)$, derived from the associated Legendre functions, which now include the ϕ -dependence:

$$Y_{lm_l}(\theta, \phi) = N_{lm_l} P_l^{m_l}(\mu) e^{im_l \phi}; \quad P_l^{m_l}(\mu) = (1 - \mu^2)^{1/2|m_l|} \frac{d^{|m_l|}}{d\mu^{|m_l|}} P_l(\mu).$$

$\Phi(\tau, \mu, \phi, \mu', \phi') = \sum_{l=0}^N \alpha_l(\tau) Y_{lm_l}(\theta, \phi)$, Where the number of terms in the expansion in l depends on the anisotropy of the phase function and the degree of accuracy required.

- b. I and Σ are expanded in Fourier cosine series in the azimuthal variable ϕ , both up to terms $m = 0, \dots, N$.

Then we have $N + 1$ equations in 2 variables, μ and τ (still integrodifferential), instead of 3 (μ, τ, ϕ):

$$\mu \frac{dI^m}{d\tau}(\tau, \mu) = I^m(\tau, \mu) - \gamma_m \int_{-1}^1 d\mu' \Phi^m(\tau, \mu, \mu') I^m(\tau, \mu') - \Sigma^m(\tau, \mu), \quad m = 0, \dots, N.$$

Sketch of the Discrete Ordinate Method

Expansion gave us a series of $N + 1$ integrodifferential equations in 2 variables.

The use of the *discrete ordinate* expansion gets rid of the integro- part to leave a system of linear differential equations.

Each of our azimuthally-independent equations (we are suppressing m -dependence for simplicity) is expanded in $\mu (= \cos \theta)$, where the most usual choice is to develop a $2-n$ stream representation with angles at the roots of the corresponding Legendre polynomials, $P_{2n}(\mu)$.

E.g., for a 2-stream expansion, $P_2 = \frac{1}{2}(3 \cos^2 \theta - 1)$; $|\mu| = 0.57735$; $|\theta| = 54.7^\circ$

$$P_4(\mu) = \frac{3}{8} \left(\frac{35\mu^4}{3} - 10\mu^2 + 1 \right)$$

4-stream: $\mu \pm 1 = 0.3400 \quad \theta \pm 1 = 70.12^\circ$

$\mu \pm 2 = 0.8611 \quad \theta \pm 2 = 30.55^\circ$

This choice is the *Gaussian quadrature* choice. (Quadrature in general means that a definite integral is being replaced by a sum: See *Wikipedia*.) Gaussian quadrature has the **marvelous** property of being *exact* for $\Phi =$ a polynomial of degree $\leq 4n$ (that is, for a $2n$ representation!) for integrated fluxes and intensities.

Expansion gives

$$\mu_{\pm i} \frac{dI}{d\tau}(\tau, \mu_{\pm i}) = I(\tau, \mu_{\pm i}) - \frac{\gamma}{2} \sum_{j=1}^n a_j \Phi(\tau, \mu_{\pm i}, \mu_j) I(\tau, \mu_j) -$$

$$\frac{\gamma}{2} \sum_{j=1}^n a_j \Phi(\tau, \mu_{\pm i}, \mu_{-j}) I(\tau, \mu_{-j}) - \Sigma(\tau, \mu_{\pm i}), \quad i = 1, n.$$

The expansion coefficients are given by the Gauss quadrature formula:

$$a_j = \frac{1}{P'_m(\mu_j)} \int_{-1}^1 \frac{P_m(\mu) d\mu}{\mu - \mu_j}, \quad \text{where } P'_m(\mu_j) = \left(\frac{dP_m}{d\mu} \right)_{\mu=\mu_j}.$$

These are tabulated extensively (see *Chandrasekhar* Chapter II and Table III), although they may now be easily computed as needed. There are other quadrature formulae, but they do not give results accurate to $\Phi \leq 4n$.

We have now replaced our integrodifferential equation in 3 variables with a set of linear differential equations which may be solved by standard methods.

Proceed by setting up a layered atmosphere with $\varepsilon, \omega, \Phi$ for each layer (interpolate from layered in z or P if necessary to layered in $\tau, \tau = \tau(\sigma)$). This adds an extra dimension (# layers) to the problem: complicated boundary value problem. It can also become complicated when τ changes rapidly.

Other complications:

1. Non-homogeneous terms (e.g., beam source);
2. Strongly-peaked Φ s may require other choice for discretization (DISORT and LIDORT discuss this)
3. Output at other than stream angles – uses a complicated (but accurate) interpolation formula *or* put in an extra stream in the calculation with zero weight (see DISORT)

and LIDORT). The most basic use is for flux and intensity integrals (see *Goody and Yung*, Chapters 2 and 8).

For a single homogeneous layer,

$$I(\tau, \mu_i) = \sum_{j=-n}^n L_j g_j(\mu_i) e^{-k_j \tau} \text{ (homogeneous)} + I_p(\tau, \mu_i) \text{ (particular solution):}$$

Solution to $2n$ first-order differential equations with constant coefficients, plus non-homogeneous terms, where the k_j and g_j are the eigenvalues and eigenvectors of the solution to the differential equations based on the discrete ordinate expansion (*cf. GY 8.30*).

The multiple-layer solution is then a complicated boundary-value problem where the intensity for each azimuthal component and stream angle must be continuous across layer interfaces.

DISORT is the standard discrete ordinate development. It is widely-used and generally available (see class website for references).

LIDORT (developed at the CfA by Rob Spurr, since founder of RT Solutions, Inc) adds calculation of the full Jacobian by a full analytical perturbation analysis of intensity field: Yields Jacobians (weighting functions) in one pass (no finite-differencing); pseudo-spherical and quasi-spherical versions available; surface BRDF; vector (polarization) version available (VLIDORT). Availability: <http://www.rtslidort.com/>.

There are *many* other approaches:

- Doubling and adding method (*e.g.*, DAK)
- Successive orders of scattering
- Monte Carlo methods
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See *Goody and Yung*, Chapter 8 for details.